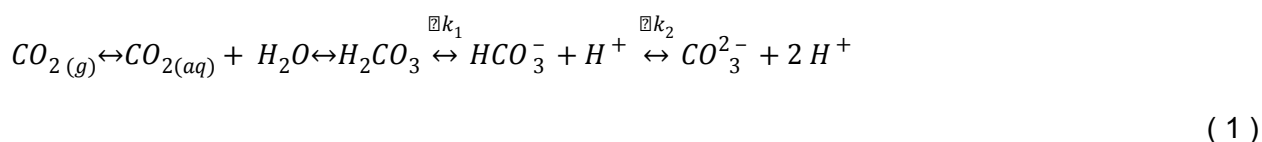


Derivation of Source Function $S(x,t)$:

The source function $S(x,t)$ quantitatively describes how the microalgae's source of inorganic C, i.e. HCO_3^- , is formed from atmospheric $CO_{2(g)}$. In a first step, Henry's law describes a thermodynamic equilibrium between gas and aqueous phase and thus determines the quantity of CO_2 that dissolves into seawater ($CO_{2(g)} \rightarrow CO_{2(aq)}$). From $CO_{2(aq)}$, carbonic acid (H_2CO_3) forms which partially dissociates in water resulting in bicarbonate (HCO_3^-) which in turn produces a small quantity of carbonate ions (CO_3^{2-}) (1). As will be discussed in the remainder, the most relevant product formed in this sequence of reactions is HCO_3^- .



Quantitating the transformation of $CO_{2(g)}$ into HCO_3^- requires a detailed analysis. First, the partial pressure of CO_2 in the atmosphere (in atm) needs to be determined:

$$pCO_2 = xCO_2 \cdot (P - pH_2O) \cdot 10^{-6} \quad (2)$$

In (2), xCO_2 represents is the mole fraction of CO_2 in the atmosphere and P the total atmospheric pressure in atm (used here: 1atm). Since, CO_2 exchange occurs at the air-water interface, saturation vapor pressure of water, pH_2O (atm) needs to be subtracted from the total pressure. Saturation water vapor pressure can be calculated by 1:

$$\ln(pH_2O) = 24.4543 - \frac{6745.09}{T} - 4.8489 \cdot \ln\left(\frac{T}{100}\right) - 0.00054 \cdot S \quad (3)$$

In (3), T represents the temperature of the air-water interface in Kelvin and S denotes the seawater salinity in g/kg. Since, CO_2 is not an ideal gas, its fugacity fCO_2 (in atm) needs to be calculated from its partial pressure and ideal gas constant R 2:

$$fCO_2 = pCO_2 \cdot \exp\left(P \cdot \frac{B + 2\delta}{R \cdot T}\right) \quad (4)$$

For calculating f^{CO_2} (4), B and δ (both in $m^3 \cdot mol^{-1}$) are required which denote the virial coefficient and the cross-virial coefficient, respectively ³:

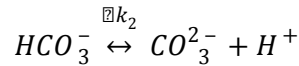
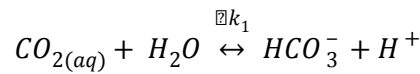
$$B = (-1636.75 + 12.0408 \cdot T - 3.27957 \cdot 10^{-2} \cdot T^2 + 3.16528 \cdot 10^{-5} \cdot T^3) \cdot 10^{-6}$$

$$\delta = (57.7 - 0.118 \cdot T) \cdot 10^{-6}$$

After calculating fugacity, the CO_2 concentration (mol/kg) dissolved in seawater can be calculated from $[CO_2] = K_0 \cdot f^{CO_2}$ with K_0 (mol/kg-atm) denoting CO_2 solubility coefficient which has been determined in ³ as:

$$K_0 = 10^{-6} \cdot \exp\left\{\frac{9345.17}{T} - 60.2409 + 23.3585 \cdot \ln\left(\frac{T}{100}\right) + S \cdot (0.023\right.$$

Dissolved $CO_{2(aq)}$ is in equilibrium with HCO_3^- and CO_3^{2-} according to:



The equilibrium constants $k_1 = \frac{[HCO_3^-] \cdot [H^+]}{[CO_2]}$ and $k_2 = \frac{[CO_3^{2-}] \cdot [H^+]}{[HCO_3^-]}$ are temperature

T ($= 296K$) and salinity $S = 35 \frac{g \text{ salt}}{kg \text{ seawater}}$ dependent functions (in mol/kg) and have been determined in ⁴ as:

$$\ln(k_1) = 2.83655 - \frac{2307.1266}{T} - 1.5529413 \cdot \ln(T) - \left(0.207608410 + \frac{4.048}{T}\right)$$

$$\ln(1 - 0.001005 \cdot S)$$

$$\ln(k_2) = -9.226508 - \frac{3351.6106}{T} - 0.2005743 \cdot \ln(T) - \left(0.1069018 + \frac{23.97}{T}\right)$$

$$\ln(1 - 0.001005 \cdot S)$$

$[H^+]$ (in mol/kg) can be determined from the pH ($= 8.20$; *buffered*) and density of the seawater

solution $\left(= 1.026 \frac{kg}{L}\right)$. With $[CO_2]$, $[H^+]$, and k_1 known, $[HCO_3^-]$ (in mol/kg) can be calculated:

$$[HCO_3^-] = \frac{k_1 \cdot [CO_2]}{[H^+]}$$

For completeness purposes, the carbonate ion concentration can be determined, too:

$$[CO_3^{2-}] = \frac{k_2 \cdot [HCO_3^-]}{[H^+]}$$

Finally, gravimetric concentration units (mol/kg) are converted to millimolar (mM) through the density of seawater (kg/L) for this modeling application.

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