

Supplemental Material

Uranium and Am isotope ratios for spiked geostandard samples.				
	$^{234}\text{U}/^{238}\text{U}$	$^{235}\text{U}/^{238}\text{U}$	$^{236}\text{U}/^{238}\text{U}$	$^{241}\text{Am}/^{243}\text{Am}$
BCR-2A	0.0171(1)	1.738(2)	0.0264(2)	0.653(2)
BCR-2B	0.0149(1)	1.517(2)	0.0230(2)	0.640(2)
BCR-2C	0.0098(1)	0.997(2)	0.0151(2)	0.606(2)
BCR-2D	0.0175(1)	1.771(2)	0.0269(2)	0.647(2)
BHVO2+	0.01043(2)	1.0003(1)	0.00152(1)	-

Plutonium isotope ratios for spiked geostandard samples.				
Sample	$^{238}\text{Pu}/^{239}\text{Pu}$	$^{240}\text{Pu}/^{239}\text{Pu}$	$^{241}\text{Pu}/^{239}\text{Pu}$	$^{242}\text{Pu}/^{239}\text{Pu}$
BCR-2x	0.00263(8)	0.2407(3)	0.00525(8)	0.01561(5)

Background-corrected isotope ratios were calculated according to:

$$R_{ab} = \frac{(N_a - N_{a,B})}{(N_b - N_{b,B})} \quad (1)$$

where N_i is the number of counts in peak i and N_{iB} is the number of background counts in peak i .

The uncertainty in isotope ratio measurements is calculated from counting statistics:

$$\sigma^2 = \sum_i \left(\frac{\delta R_{ab}}{\delta N_i} \right)^2 \sigma_i^2 \quad (2)$$

where σ_i is the uncertainty in N_i . Applying (2) to (1) and recognizing that for counting statistics $\sigma_i = \sqrt{N_i}$ gives:

$$\sigma^2 = \frac{N_a}{(N_b - N_{b,B})^2} + \frac{N_{a,B}}{(N_b - N_{b,B})^2} + \frac{N_b(N_a - N_{a,B})^2}{(N_b - N_{b,B})^4} + \frac{N_{b,B}(N_a - N_{a,B})}{(N_b - N_{b,B})^4} \quad (3)$$

which, after simplification, yields

$$\sigma = R_{ab} \sqrt{\frac{N_a + N_{a,B}}{(N_a - N_{a,B})^2} + \frac{N_b + N_{b,B}}{(N_b - N_{b,B})^2}} \quad (4)$$

