

Table ESI 2: ID-TIMS data

Sample	Compositional Parameters							Radiogenic Isotope Ratios							Isotopic ages, Ma										
	Wt.	U	Th	Pb	Pb*	Pb _c	²⁰⁶ Pb	²⁰⁶ Pb	²⁰⁷ Pb	2 sigma	²⁰⁷ Pb	2 sigma	²⁰⁶ Pb	2 sigma	corr.	²⁰⁷ Pb	2 sigma	²⁰⁷ Pb	2 sigma	²⁰⁶ Pb	2 sigma	²⁰⁶ Pb	2 sigma		
	mg	ppm	U	ppm	Pb _c	(pg)	²⁰⁴ Pb	²⁰⁶ Pb	²⁰⁶ Pb	% err	²³⁵ U	% err	²³⁸ U	% err	coef.	²⁰⁶ Pb	±	²³⁵ U	±	²³⁸ U	±	²⁰⁶ Pb	±	²³⁸ U	±
(a)	(b)	(c)	(d)	(c)	(e)	(f)	(g)	(g)	(h)	(g)	(h)	(g)	(h)	(g)	(h)	(i)	(h)	(i)	(h)	(i, j)	(h)	(l, k)	(l)	(i)	
KPT-04																									
KPT014-13-1	0.0084	937	0.167	0.40	0.04	3.19	21.467	0.102	0.046144	44.1	0.000216	47.2	0.0000340	8.59	0.44	5.13971	847	0.22	0.10	0.219	0.010	0.203	0.015		
KPT014-13-2	0.0069	1119	0.432	0.13	0.19	0.78	29.065	0.264	0.046053	31.5	0.000202	33.6	0.0000317	2.86	0.75	0.38683	606	0.20	0.07	0.205	0.003	0.187	0.011		
KPT014-13-3	0.0084	933	0.398	0.25	0.10	1.92	24.394	0.216	0.046090	35.3	0.000237	37.7	0.0000373	3.99	0.63	2.35379	703	0.24	0.09	0.241	0.006	0.229	0.015		
KPT014-13-6	0.0077	557	0.859	0.25	0.16	1.66	27.094	0.348	0.046145	41.0	0.000412	42.5	0.0000647	2.62	0.58	5.21123	899	0.42	0.18	0.417	0.009	0.410	0.027		
KPT014-13-8	0.0065	1032	0.551	0.28	0.11	1.62	24.779	0.287	0.046054	39.8	0.000241	42.0	0.0000379	3.57	0.65	0.45121	801	0.24	0.10	0.244	0.005	0.230	0.019		
KPT014-13-7	0.0058	717	0.370	0.24	0.10	1.24	24.374	0.186	0.046054	65.6	0.000269	68.1	0.0000424	4.10	0.63	0.44763	1334	0.27	0.19	0.273	0.007	0.262	0.016		
KPT014-13-10	0.0058	1693	0.586	0.28	0.15	1.41	26.631	0.339	0.046112	33.1	0.000206	34.9	0.0000325	2.83	0.66	3.48032	646	0.21	0.07	0.209	0.003	0.190	0.017		

(a) z1, z2 etc. are labels for fractions composed of single zircon grains or fragments; all fractions annealed and chemically abraded after Mattinson (2005).

(b) Nominal fraction weights measured after chemical abrasion.

(c) Nominal U and total Pb concentrations subject to uncertainty in weighting zircons.

(d) Model Th/U ratio calculated from radiogenic ²⁰⁶Pb/²⁰⁶Pb ratio and ²⁰⁷Pb/²³⁵U age.

(e) Pb* and Pb_c represent radiogenic and common Pb, respectively; mol % ²⁰⁶Pb* with respect to radiogenic, blank and initial common Pb.

(f) Measured ratio corrected for spike and fractionation only. Mass fractionation correction of 0.11 ± 0.02 (1-sigma) %/amu (atomic mass unit) was applied to all single-collector Daly analyses, based on analysis of NBS-981 and NBS-982.

(g) Corrected for fractionation, spike, and common Pb; all common Pb was assumed to be procedural blank: ²⁰⁶Pb/²⁰⁴Pb = 18.30 ± 0.26%; ²⁰⁷Pb/²⁰⁴Pb = 15.47 ± 0.32%; ²⁰⁸Pb/²⁰⁴Pb = 37.60 ± 0.74% (all uncertainties 1-sigma). ²⁰⁶Pb/²³⁸U and ²⁰⁷Pb/²³⁵U ratios corrected for initial disequilibrium in ²³⁰Th/²³⁸U using Th/U [magma] = 3.3 (KPT04),

(h) Errors are 2-sigma, propagated using the algorithms of Schmitz and Schoene (2007) and Crowley et al. (2007).

(i) Calculations are based on the decay constants of Jaffey et al. (1971). ²⁰⁶Pb/²³⁸U and ²⁰⁷Pb/²⁰⁶Pb ages corrected for initial disequilibrium in ²³⁰Th/²³⁸U using Th/U [magma]

(j) Disequilibrium U-Th corrected after Schärer, 1984 .

(k) Disequilibrium U-Th corrected after Sakata *et al.*, 2013 .

Derivation of the advance Th disequilibrium correction

The distribution ratio D of U and Th between source magma and zircon is defined as follows:

$$D_{Zircon/Source\ Magma}^U = \frac{U_{Zircon}^{Initial}}{U_{Source\ Magma}^{Initial}} \quad (ESI\ 1)$$

$$D_{Zircon/Source\ Magma}^{Th} = \frac{Th_{Zircon}^{Initial}}{Th_{Source\ Magma}^{Initial}} \quad (ESI\ 2)$$

where initial represents the time of zircon formation. Now, as $f_{Th/U}$ represent the ratio of the D of U and Th, respectively between source magma and zircon crystal, the f is defined as follows:

$$f_{Th/U} = \frac{D_{Zircon/SourceMagma}^{Th}}{D_{Zircon/SourceMagma}^U} \quad (ESI\ 3)$$

According to Schärer (1984), by means of using $f_{Th/U}$, it is possible to calculate the time variation of $^{206}Pb/^{238}U$ ratio considering the effect of initial disequilibrium. This relationship is expressed with following equation:

$$\frac{^{206}Pb}{^{238}U} = (e^{\lambda_{238} t} - 1) + \frac{\lambda_{238}}{\lambda_{230}} (f_{Th/U} - 1) \quad (1)$$

However, although the growth of $^{206}Pb/^{238}U$ ratio can still be calculated after ^{230}Th reaches radioactive equilibrium using equation (1), this equation cannot do accurate correction in the scope of time before reaching equilibrium (up until about 0.3 Ma for ^{230}Th).

In fact, the isotopic growth curve defined from equation (1) does not pass the origin point when $t = 0$ ($^{206}Pb^* = 0$). Therefore, for the chronology of the zircon crystals with the age of less than 1 Ma, a more accurate correction has to be applied.

We assume the following conditions here:

- Radioactive equilibrium of the uranium decay series has been established in the source magma from which the zircon crystallizes.
- $^{234}U/^{238}U$ ratio is the same in zircons and source magma (^{234}U and ^{238}U hold the radioactive equilibrium).
- Initial lead content (taken to zircon crystal from source magma at the time of zircon formation) can be regarded ($Pb^*_{initial} = 0$).

Because the radioactive equilibrium is established in the source magma (assumption (a)), the following relationship holds:

$${}^{238}\text{U}_{\text{SourceMagma}}^{\text{Initial}} \lambda_{238} = {}^{234}\text{U}_{\text{SourceMagma}}^{\text{Initial}} \lambda_{234} = {}^{230}\text{Th}_{\text{SourceMagma}}^{\text{Initial}} \lambda_{230} \quad (\text{ESI 4})$$

In addition, considering assumption (b) and (c), relationship about ${}^{238}\text{U}$, ${}^{234}\text{U}$, and ${}^{206}\text{Pb}$ in zircon crystal is expressed as follows:

$${}^{238}\text{U}_{\text{Zircon}}^{\text{Initial}} \lambda_{238} = {}^{234}\text{U}_{\text{Zircon}}^{\text{Initial}} \lambda_{234} \quad (\text{ESI 5})$$

$${}^{206}\text{Pb}_{\text{Zircon}}^{\text{Initial}} = 0 \quad (\text{ESI 6})$$

The following relationship is established in radiogenic ${}^{206}\text{Pb}$ and ${}^{230}\text{Th}$ at the age t .

$$\frac{d {}^{230}\text{Th}_{\text{Zircon}}}{dt} = {}^{234}\text{U}_{\text{Zircon}} \lambda_{234} - {}^{230}\text{Th}_{\text{Zircon}} \lambda_{230} \quad (\text{ESI 7})$$

$$\frac{d {}^{206}\text{Pb}_{\text{Zircon}}^*}{dt} = {}^{230}\text{Th}_{\text{Zircon}} \lambda_{230} \quad (\text{ESI 8})$$

After solving equation (ESI 1) – (ESI 8) and using the approximation of $\lambda_{230} - \lambda_{238} \cong \lambda_{230}$, we can finally get the following equation that represents the relationship between the ${}^{206}\text{Pb}/{}^{238}\text{U}$ ratio and age t .

$$\frac{{}^{206}\text{Pb}_{\text{Zircon}}^*}{{}^{238}\text{U}_{\text{Zircon}}} = (e^{\lambda_{238} t} - 1) + \frac{\lambda_{238}}{\lambda_{230}} \left(f_{\text{Th}} - 1 \right) (1 - e^{-\lambda_{230} t}) e^{\lambda_{238} t} \quad (3)$$

Figure ESI 1: Evaluation of common Pb on example KPT04-24 (17. April 2013):

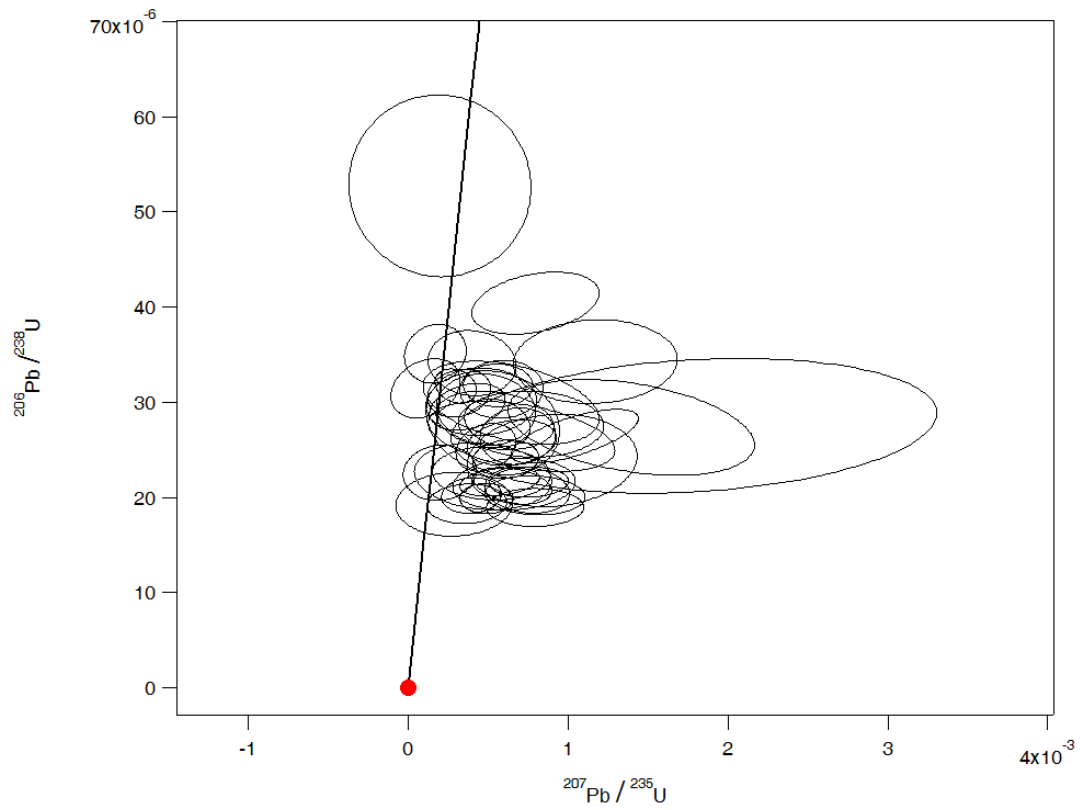


Figure ESI 2: Evaluation of common Pb on example KPT04-36 (17. April 2013) Inherited core not shown:

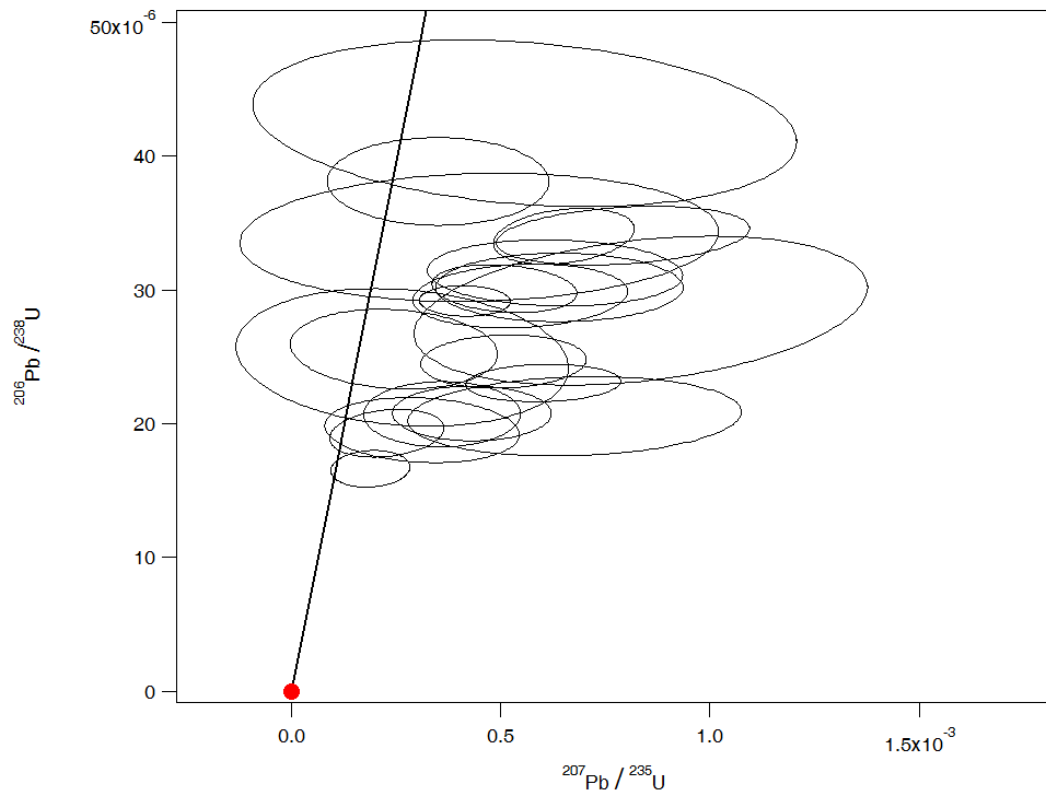


Figure ESI 3: Evaluation of common Pb on example KPT04-24 (01. February 2013):

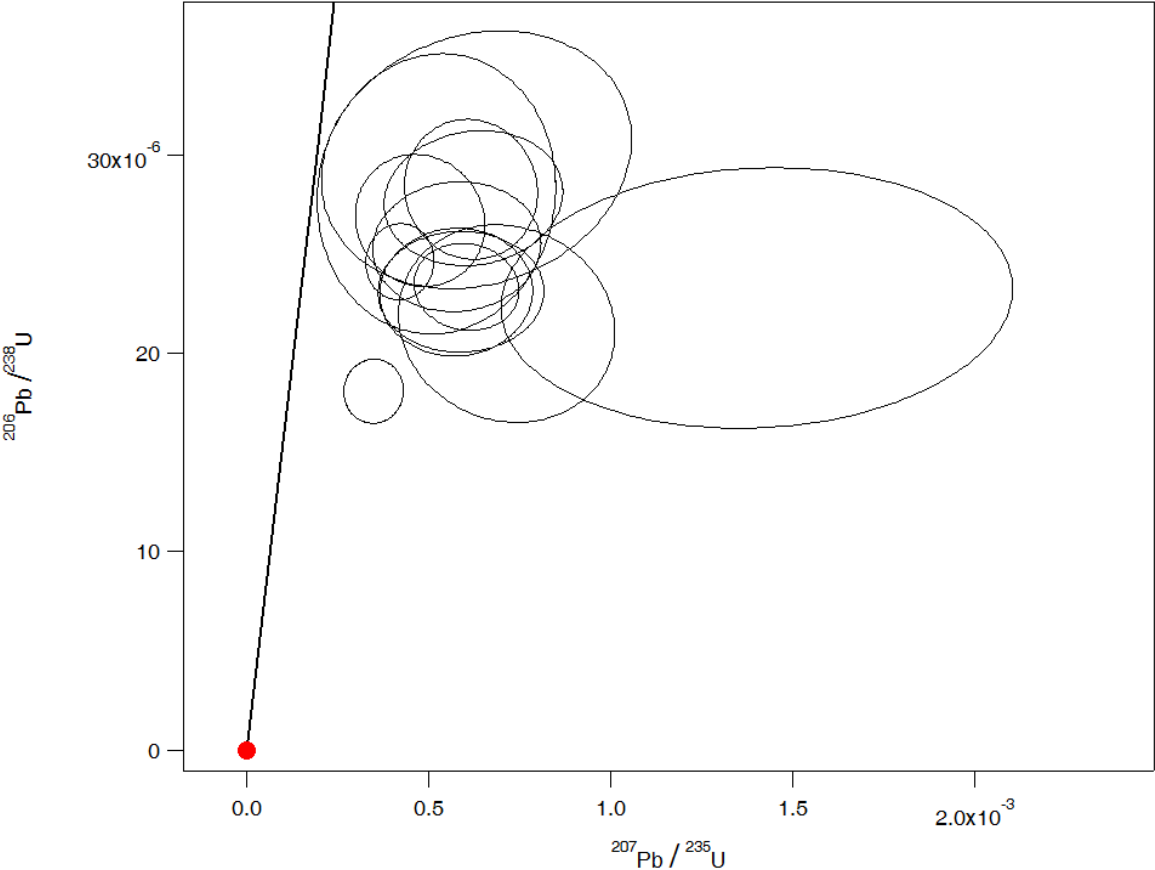


Figure ESI 4: Evaluation of common Pb on example KPT04-36 (01. February 2013) Inherited core not shown:

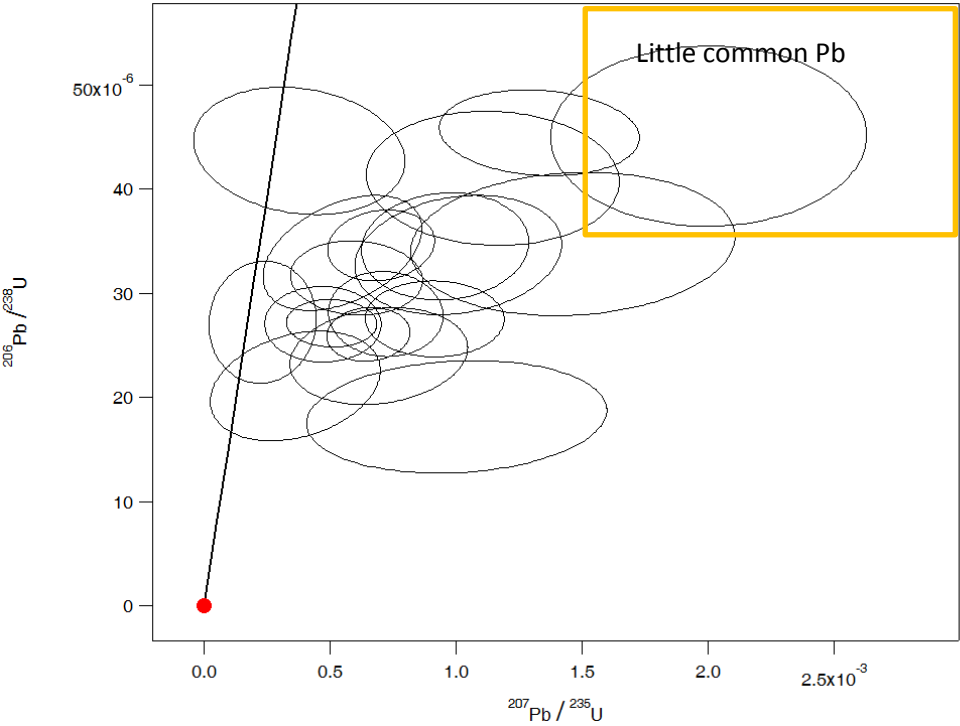


Figure ESI 5: Evaluation of common Pb on example NS07:

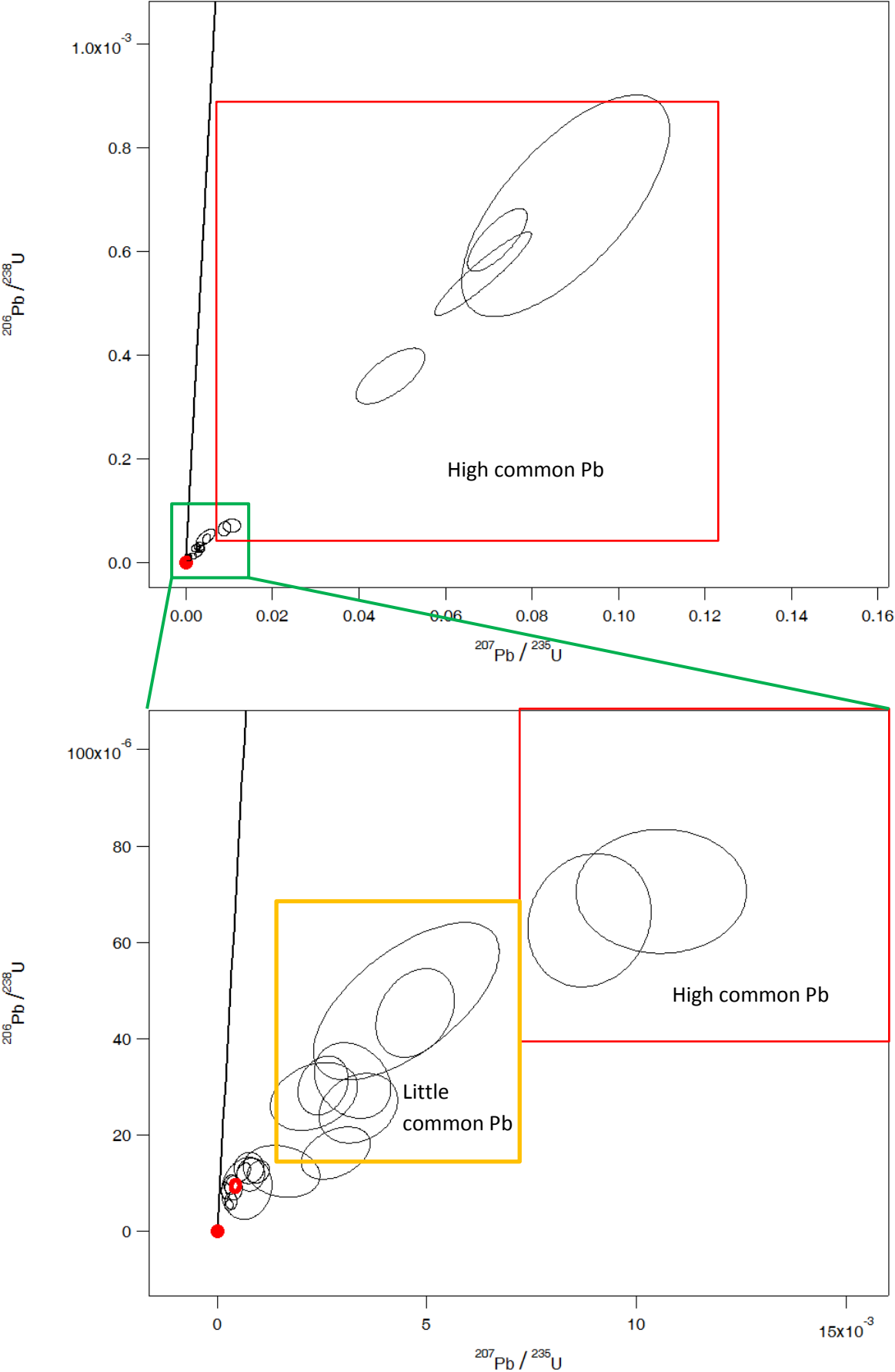


Figure ESI 6: Evaluation of common Pb on example NS24:

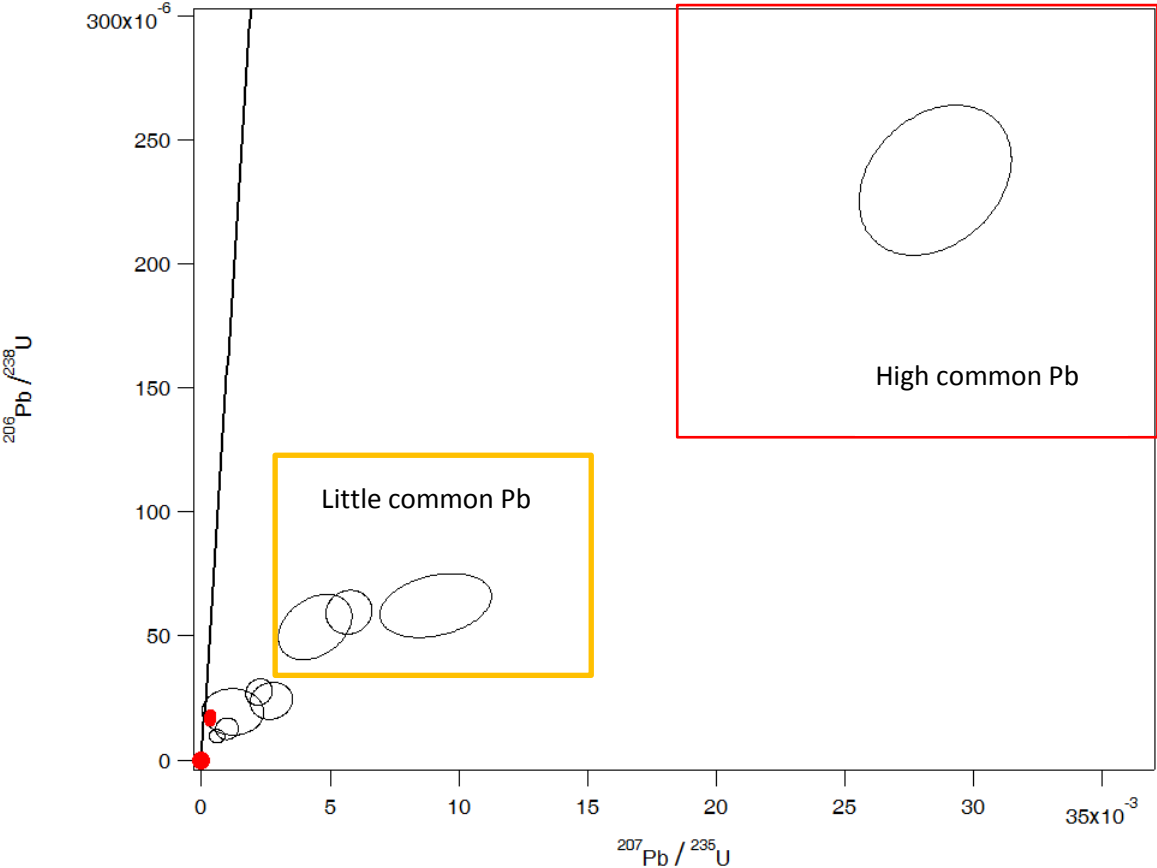


Figure ESI 7: Evaluation of common Pb on example NISP2(1):

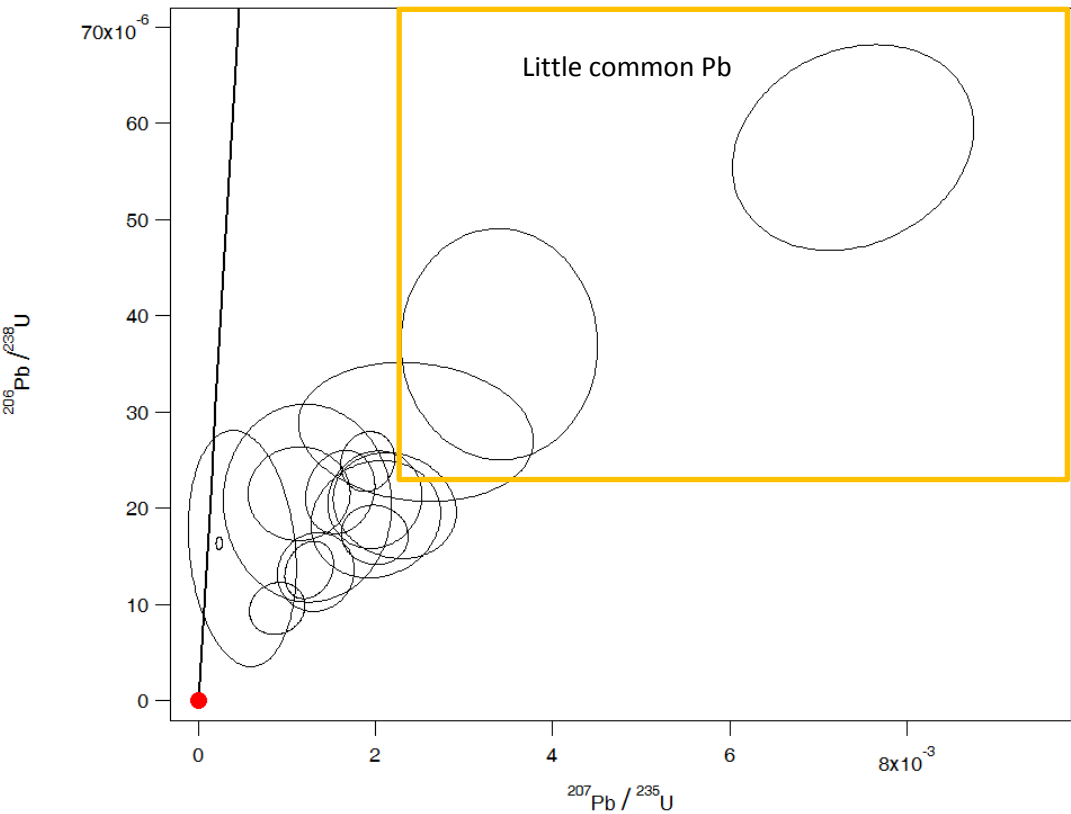


Figure ESI 8 Probability plot of all KOS ages (SHRIMP, ID-TIMS and LA-ICP-MS) does not show a normal distribution.

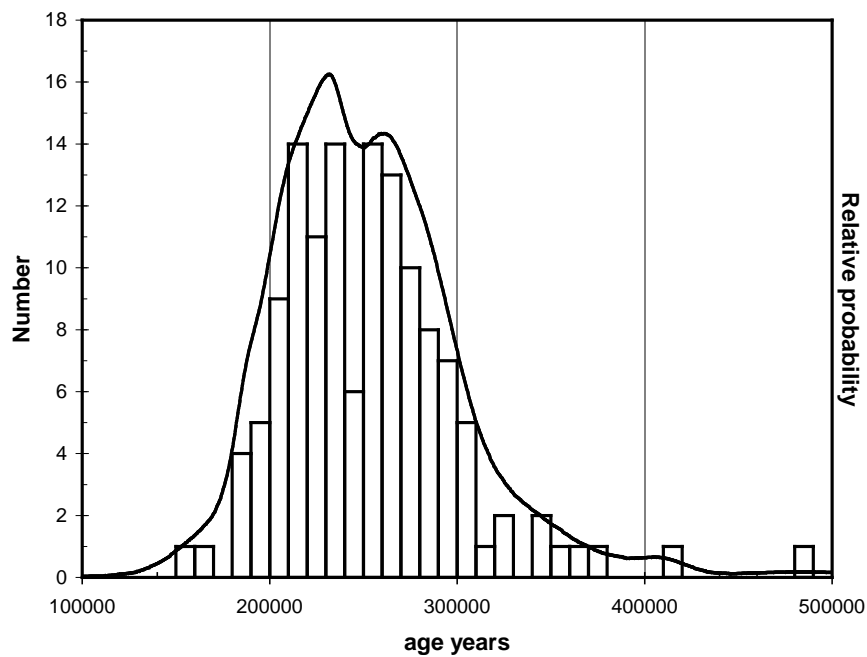


Figure ESI 9: All ages for the KOS sample together showing continuous crystallization over 140ka. The calculation of a mean age seems meaningless.

