



**ANALYSIS OF NATURAL CONVECTION  
IN A FLUID HEATED FROM BELOW  
(THE RAYLEIGH-BÉNARD PROBLEM)**

**CHEN 304  
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# NON-ISOTHERMAL SYSTEMS

Example: Natural Convection

Non-uniform temperature causes density change within volume

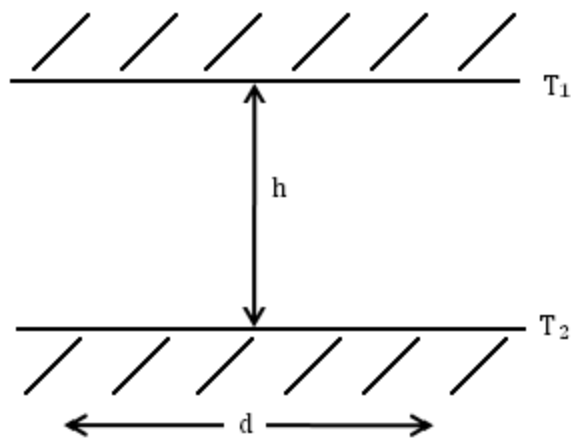
- → Can create unstable system
- → Flow

Oceans, Atmosphere, Earth's core  
→ Lava Lamps



# RAYLEIGH-BÉNARD PROBLEM

- Fluid confined between two horizontal surfaces at different temperatures

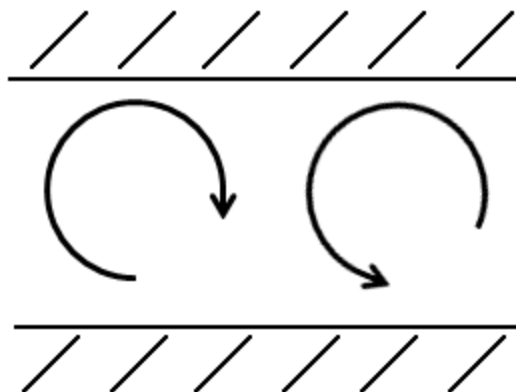


If  $T_2 > T_1$ , fluid is less dense near bottom  
→ creates a “top heavy” arrangement → unstable



# RAYLEIGH-BÉNARD PROBLEM

- Under the right conditions, a continuous circulatory flow can be generated



Convection cells with characteristic size and spacing

Conditions for flow should depend on

- 1) Geometry ( $h$ ,  $d$ )
- 2) Fluid properties
- 3) Temperature difference between top and bottom



## HOW TO ANALYZE?



- Conservation of Mass and Momentum
  - (Navier-Stokes)
- Conservation of Energy
  - (accounts for thermal effects, CHEN 323)



## HOW TO ANALYZE?



- Energy equation adds another system of partial differential equations (PDE) that are coupled with the PDE's in Navier-Stokes
- Very difficult to solve
- Must use Computational Fluid Dynamics CFD



# CONSERVATION EQUATIONS

## Conservation of Mass

$$\nabla \cdot v = 0$$

## Conservation of Momentum (Navier-Stokes)

$$\rho \left( \frac{\partial v}{\partial t} + v \cdot \nabla v \right) = -\nabla P + \rho g + \mu \nabla^2 v$$

## Conservation of Energy (CHEN 323)

$$\rho C_p \left( \frac{\partial T}{\partial t} + v \cdot \nabla T \right) = -k \nabla^2 T$$

- Equations are coupled through velocity
- Also linked by temperature
- Temperature differences are the driving force for fluid motion



## ANALYZE

- First step: think about body force term
- Changes in density with respect to the base state causes variable in hydrostatic pressure that can become unstable (“top heavy”)
- Base state
  - No fluid motion ( $v = 0$ )
  - $\rho = \rho_0$





## ANALYZE

- Substitute base state into Navier-Stokes

### Hydrostatic Equation

$$0 = -\nabla P_h + \rho_o g$$

Thus we can express

$$P = P_h + P_d$$

**“hydrostatic” pressure**  
that would exist if  $\rho = \rho_o$   
everywhere

**“dynamic” pressure**  
due to density variation in  
temperature gradient



## ANALYZE

$$\nabla P = \nabla P_h + \nabla P_d$$

$$\nabla P_h = \rho_o g$$

$$\begin{aligned}\rho \left( \frac{\partial v}{\partial t} + v \cdot \nabla v \right) &= -(\nabla P_h + \nabla P_d) + \rho g + \mu \nabla^2 v \\ &= -(\rho_o g + \nabla P_d) + \rho g + \mu \nabla^2 v \\ &= -\nabla P_d + (\rho - \rho_o)g + \mu \nabla^2 v\end{aligned}$$

- Now, we know that changing  $T$  also changes  $\rho$  (and other fluid properties).
- How do we express mathematically?



## BOUSSINESQ APPROXIMATION

- 1) Variations in fluid properties with temperature are neglected, except for density
- 2) Temperature dependence of density is only important in buoyancy term
- 3) Here, density changes linearly with  $T$

Valid for small temperature variations

$$\rho = \rho_o [1 - \beta(T - T_o)]$$

At base state

$\beta$  - thermal expansion coefficient (fluid property)



## BOUSSINESQ APPROXIMATION

- Substitute into N-S

$$\rho \left( \frac{\partial v}{\partial t} + v \cdot \nabla v \right) = -\nabla P_d + \{[\cancel{\rho_o} - \rho_o \beta(T - T_o)] - \cancel{\rho_o}\}g + \mu \nabla^2 v$$

$$\rho \left( \frac{\partial v}{\partial t} + v \cdot \nabla v \right) = -\nabla P_d + \rho_o \beta(T - T_o)g + \mu \nabla^2 v$$



## NON-DIMENSIONALIZE

$$v^* = \frac{v}{V}$$

$$t^* = \frac{t}{L/V}$$

$$T^* = \frac{T - T_o}{T_2 - T_o} = \theta$$

$$\nabla^* = L\nabla$$

$$P^* = \frac{P}{\left(\frac{\mu V}{L}\right)}$$

$$\nabla^{*2} = L^2\nabla^2$$



## SUBSTITUTE

$$\rho_o \frac{V^2}{L} \left( \frac{\partial v^*}{\partial t^*} + v^* \cdot \nabla^* v^* \right) = - \frac{\mu V}{L^2} \nabla^* P^* - \rho_o \beta \underbrace{([\theta(T_2 - T_o) + \cancel{T_o}] - \cancel{T_o})}_{T} g + \frac{\mu V}{L^2} \nabla^{*2} v^*$$

$$\left( \frac{\partial v^*}{\partial t^*} + v^* \cdot \nabla^* v^* \right) = - \underbrace{\frac{\mu}{\rho_o V L}}_{\frac{1}{\text{Re}}} \nabla^* P^* - \underbrace{\frac{\cancel{\rho_o} \beta (T_2 - T_o) L}{\cancel{\rho_o} V^2}}_{\text{Buoyancy term}} \theta g + \underbrace{\frac{\mu}{\rho_o V L}}_{\frac{1}{\text{Re}}} \nabla^{*2} v^*$$



## BUOYANCY

- Let's look at the buoyancy term
- If gravity acts in the  $-z$  direction

$$\mathbf{g} = -g \hat{\mathbf{e}}_z$$

$$\text{Buoyancy term} = + \frac{g\beta(T_2 - T_o)L}{V^2} \theta \hat{\mathbf{e}}_z$$



# FLUID PROPERTIES

Can we identify a characteristic velocity  $V$ ?

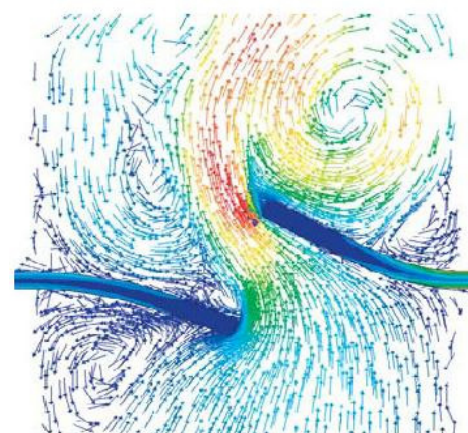
$$V = \frac{L}{T}$$

Choose timescale of momentum diffusion

$$T = \frac{L^2}{\nu}$$

$\frac{\mu}{\rho}$

Kinematic viscosity  
units =  $L^2/T$





## SUBSTITUTE

$$\text{Buoyancy term} = + \frac{g\beta(T_2 - T_o)L^3}{\nu^2} \theta \hat{e}_z$$

This gives us the following dimensionless equation:

$$\left( \frac{\partial v^*}{\partial t^*} + v^* \cdot \nabla^* v^* \right) = -\frac{1}{\text{Re}} \nabla^* P^* + \text{Gr} \theta \hat{e}_z + \frac{1}{\text{Re}} \nabla^{*2} v^*$$

where

$$\text{Gr} = \frac{g\beta(T_2 - T_o)L^3}{\nu^2}$$

Grashof number



- Multiply both sides of the equation by Prandtl number

$$\text{Pr} = \frac{\nu}{\alpha} = \frac{\text{momentum diffusion}}{\text{thermal diffusion}}$$

$$\alpha = \frac{k}{\rho C_p}$$

This gives a group

$$\text{GrPr} = \text{Ra} = \frac{g\beta(T_2 - T_o)L^3}{\nu \alpha}$$

**Rayleigh  
number**



- $Gr$ ,  $Ra$  express the ratio of the destabilizing forces
  - $(\Delta T, g, h)$
- Create density differences opposed by the restoring forces
  - viscous and thermal diffusion that act to smooth out these differences



## EUREKA



- Makes sense!
- Increasing  $g$ ,  $\Delta T$ ,  $h$  makes it unstable
- Increasing  $\nu$ ,  $\alpha$  makes it more stable
- This shows us how the basic parameters influence the flow



# ENERGY BALANCE

- Non-dimensionalize the energy balance

$$\rho C_p \left\{ \frac{V}{L} \frac{\partial}{\partial t^*} [\theta(T_2 - T_o) + T_o] + \frac{V}{L} \nabla^* \cdot v^* [\theta(T_2 - T_o) + T_o] \right\} \\ = \frac{k}{L^2} \nabla^{*2} [\theta(T_2 - T_o) + T_o]$$

$$\left( \frac{\partial \theta}{\partial t^*} + v^* \cdot \nabla^* \theta \right) = \frac{k}{LV\rho C_p} \nabla^{*2} \theta$$

multiply by  $\frac{\mu}{\mu}$

$$\left( \frac{\partial \theta}{\partial t^*} + v^* \cdot \nabla^* \theta \right) = \frac{\mu}{LV\rho} \cdot \frac{k}{\mu C_p} \nabla^{*2} \theta$$

$$\left( \frac{\partial \theta}{\partial t^*} + v^* \cdot \nabla^* \theta \right) = \frac{1}{\text{Re}} \cdot \frac{1}{\text{Pr}} \nabla^{*2} \theta$$



## EQUATIONS

- Now we have a system of 3 coupled PDE's to solve

$$\nabla^* \cdot v^* = 0$$

$$\left( \frac{\partial v^*}{\partial t^*} + v^* \cdot \nabla^* v^* \right) = -\frac{1}{\text{Re}} \nabla^* P^* + \text{Gr} \theta \hat{e}_z + \frac{1}{\text{Re}} \nabla^{*2} v^*$$

$$\frac{\partial \theta}{\partial t^*} + v^* \cdot \nabla^* \theta = \frac{1}{\text{Re}} \cdot \frac{1}{\text{Pr}} \nabla^{*2} \theta$$

where

$$\text{Re} = \frac{\rho V L}{\mu}$$

$$\text{Pr} = \frac{\nu}{\alpha} = \frac{\mu C_p}{k}$$

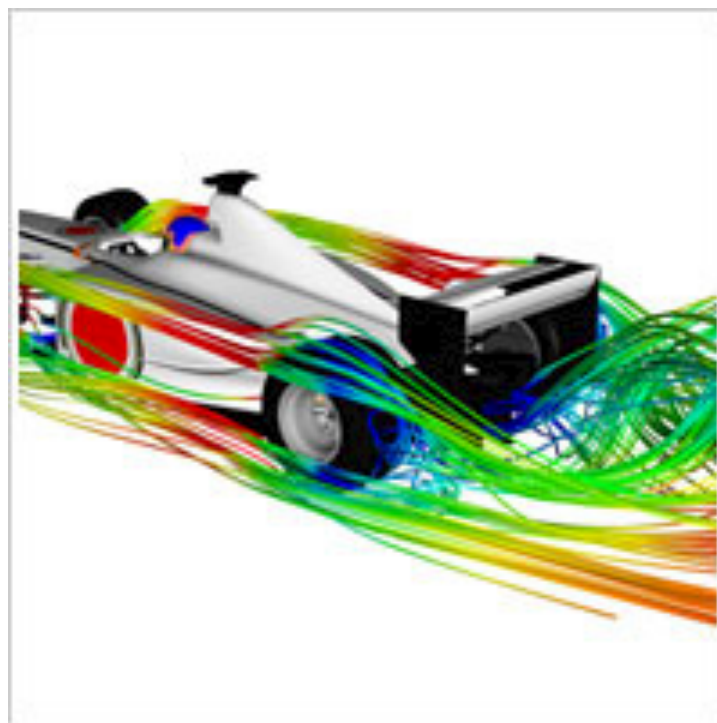
$$\text{Gr} = \frac{g \beta (T_2 - T_o) L^3}{\nu^2}$$

**Additionally**

Initial conditions and  
boundary conditions  
for velocity and  
temperature



# SOLVE?



## Computational Fluid dynamics (CFD)

