

## Supplemental Material

### Stabilities of Thiomolybdate Complexes of Iron; Implications for Retention of Essential Trace Elements (Fe, Cu, Mo) in Sulfidic Waters.

George R. Helz<sup>1\*</sup>, Britt E. Erickson<sup>1,2</sup>, Trent P. Vorlicek<sup>3</sup>

- <sup>1.</sup> Department of Chemistry and Biochemistry, University of Maryland, College Park MD 20742
- <sup>2.</sup> American Chemical Society, 1155 16<sup>th</sup> St, NW, Washington DC 20036.
- <sup>3.</sup> Department of Chemistry and Geology, Minnesota State University, Mankato MN 56001

\* To whom correspondence should be addressed: [helz@umd.edu](mailto:helz@umd.edu), 301-405-1797

## I. Details Concerning Experimental Methods

**Synthesis of Pyrrhotite.** Synthetic pyrrhotite was prepared according to the method of Kullerud and Yoder (1959) from equimolar amounts of iron powder (99.999% pure, Johnson Matthey) and elemental sulfur powder (99.998%, Aldrich Gold Label). Prior to mixing with the sulfur, the iron powder was heated to 500 °C under a continuous stream of 95% Ar–5% H<sub>2</sub> for two hours in order to reduce any oxides from the surface of the iron. The iron was allowed to cool to room temperature while still under Ar/H<sub>2</sub> and was then immediately mixed with the sulfur. The iron-sulfur mixture was finely ground with a mortar and pestle, transferred into a quartz tube, and fusion sealed under vacuum. The contents in the evacuated quartz tube were allowed to react at room temperature for one day. The tube was then heated in a furnace, increasing the temperature 50 °C per day, to a final temperature of 550 °C. After holding at 550 °C for one week, the mixture was cooled to room temperature, and was finely ground to increase homogeneity. The solid was then reheated to 550 °C in a new evacuated quartz tube for an additional three days. In the final product, the atomic Fe deficiency that is characteristic of pyrrhotite was not greater than 2% based on the  $d_{102}$  spacing (Arnold, 1966). Within uncertainty, the material was stoichiometric, hexagonal Fe<sub>1.00</sub>S (troilite).

**Reagents.** Solutions were prepared from N<sub>2</sub>-purged deionized water. The N<sub>2</sub> was passed through a Ridox oxygen absorbing column prior to entering solutions. Sulfide stock solutions were prepared by bubbling H<sub>2</sub>S gas through deoxygenated NaOH. This approach minimizes the polysulfide contamination found in Na<sub>2</sub>S•9H<sub>2</sub>O reagents. All Mo-containing solutions were prepared from (NH<sub>4</sub>)<sub>6</sub>Mo<sub>7</sub>O<sub>24</sub>•4H<sub>2</sub>O reagent. Glassware was soaked overnight in 5% nitric acid to remove trace metals before rinsing with deionized water.

Thiomolybdate solutions of varying pH,  $\Sigma S^{2-}$ ,  $\Sigma Mo$ , and  $\Sigma Cl^-$  were prepared from two 0.01 M  $\Sigma Mo/0.1 M \Sigma S^{2-}$  stock solutions that had been allowed to react for a week beforehand. One stock solution contained no NH<sub>3</sub>/NH<sub>4</sub><sup>+</sup> beyond that in the Mo reagent whereas the other was buffered with an additional 0.19 M NH<sub>3</sub>/NH<sub>4</sub><sup>+</sup>. In both stock solutions, MoS<sub>4</sub><sup>2-</sup> was the most stable form of Mo<sup>VI</sup> because of the high

sulfide concentration. Based on UV-vis spectra,  $\text{MoS}_4^{2-}$  indeed predominated in the high- $\text{NH}_4$  stock solution, but owing to slower sulfidation kinetics at low  $\text{NH}_4^+$ ,  $\text{MoOS}_3^{2-}$  and  $\text{MoO}_2\text{S}_2^{2-}$  predominated in the other stock solution at the time when the samples were prepared. Sulfidation $\leftrightarrow$ hydrolysis equilibrium among thiomolybdates is facilitated by  $\text{NH}_4^+$ , which acts as an acid catalyst (Erickson and Helz, 2000).

**Sulfide gains and losses.** In most but not all samples, dissolved  $\Sigma\text{S}^{\text{II}}$  in the final solutions differed relative to blank solutions containing no pyrrhotite. The  $\Sigma\text{S}^{\text{II}}$  differences were biased in the direction of losses from the pyrrhotite-containing solutions by an average of 6%, but in a few samples there were apparent gains. The losses, which did not vary systematically with solution composition, were generally too large to be explained by reduction of  $\text{Mo}^{\text{VI}}$  to  $\text{Mo}^{\text{V}}$  or to  $\text{Mo}^{\text{IV}}$ , a process that would transform  $\text{S}^{\text{II}}$  to  $\text{S}^0$  in polysulfides. We are uncertain of the explanation for the losses, but  $\text{H}_2\text{S}$  off-gassing during sample handling may have contributed. A procedural artifact of this kind is indicated by the fact that batches of samples processed at the same time tended to have similar, though not identical  $\Sigma\text{S}^{\text{II}}$  losses. The data that we report below are the measured  $\Sigma\text{S}^{\text{II}}$  in the pyrrhotite-equilibrated solutions.

**Procedures.** Individual test samples were prepared in a  $\text{N}_2$ -filled glove box. Aliquots of Mo stock solutions were combined with aliquots of a freshly prepared 1.00 M NaHS stock solution to give desired concentrations of Mo and  $\text{HS}^-$ . Sufficient NaCl was added to give desired  $\text{Cl}^-$  concentrations, and the solutions were brought to volume with deoxygenated, deionized water. In some solutions, the pH was adjusted with 2 M HCl and/or 3 M NaOH to give the desired initial pH.

Approximately 35 mL aliquots of each prepared thiomolybdate solution were transferred into glass ampoules containing approximately 50 mg of pyrrhotite. The ampoules were capped with rubber balloons, removed from the  $\text{N}_2$  glove box, and immediately sealed using a propane/ $\text{O}_2$  flame. Sealed samples were allowed to react at room temperature (23-27 °C) with periodic shaking by hand to aid mixing. Reaction times in the ampoules ranged from 22 to 264 days.

After the reaction time, ampoules were opened in a N<sub>2</sub> filled glove box. Solutions were filtered (0.02 µm pore diameter, Whatman Anotop 25), and the pH and total sulfide immediately determined. In some cases, a UV-visible absorption spectrum was obtained. Residual pyrrhotite was washed with deionized water and methanol, and dried by vacuum filtration in the glove box. Selected samples were examined by X-ray diffraction, confirming that the pyrrhotite composition did not change during the experiments.

**Analyses.** Sulfide was determined by potentiometric (Ag/Ag<sub>2</sub>S electrode) titration with HgCl<sub>2</sub> after raising the sample pH to 13 with NaOH. This titration determines not only free sulfide (H<sub>2</sub>S+HS<sup>-</sup>), but also sulfide bound to thiomolybdates and polysulfides, if any. Titrations were performed on an automatic titrator and have an estimated precision of ±5%. The pH was determined by glass electrode against a double junction reference electrode. Total molybdenum was determined by flame (N<sub>2</sub>O-acetylene) atomic absorption at a wavelength of 313.3 nm. Samples for Mo determination were diluted volumetrically to give ΣMo = 3–5 ppm. Sulfide was oxidized to sulfate with BrCl and then 1000 ppm Al was added to suppress interference from Fe (David, 1968). The Al content of samples and standards were matched. Total iron was determined using flame (air-acetylene) atomic absorption at 248.3 nm. In only a few cases, dilution with deionized water was necessary to keep sample concentrations within the optimum working range. Precision in both Mo and Fe determinations is estimated to be ±5%.

## II. Solubility Data.

Table S-1. Solubility data

| Run No. | pH   | ΣS <sup>-II</sup> (M) | ΣMo (M)  | ΣFe (M)  | I (M) | ΣNH <sub>3</sub> (M) | Reaction time (d) |
|---------|------|-----------------------|----------|----------|-------|----------------------|-------------------|
| 5.1     | 8.73 | 3.00E-03              | 3.84E-04 | 1.00E-05 | 0.71  | 6.24E-03             | 24                |
| 5.2     | 8.73 | 4.80E-03              | 3.84E-04 | 1.50E-05 | 0.71  | 6.24E-03             | 28                |
| 5.3     | 8.82 | 6.50E-03              | 4.12E-04 | 7.00E-06 | 0.71  | 6.00E-03             | 28                |
| 5.4     | 8.85 | 8.40E-03              | 3.84E-04 | 1.80E-05 | 0.71  | 5.91E-03             | 28                |
| 12.1    | 9.08 | 3.86E-02              | 4.03E-04 | 5.00E-06 | 0.05  | 4.62E-03             | 26                |
| 12.2    | 9.00 | 2.06E-02              | 4.03E-04 | 3.00E-06 | 0.03  | 4.94E-03             | 26                |
| 12.3    | 8.92 | 1.15E-02              | 4.12E-04 | 1.00E-06 | 0.02  | 5.25E-03             | 26                |
| 9.1     | 9.15 | 1.22E-02              | 4.13E-04 | 3.60E-05 | 1.41  | 5.21E-03             | 22                |
| 9.2     | 9.14 | 1.25E-02              | 4.13E-04 | 1.80E-05 | 0.71  | 4.91E-03             | 22                |

|      |      |          |          |          |      |          |     |
|------|------|----------|----------|----------|------|----------|-----|
| 9.3  | 9.14 | 1.23E-02 | 4.13E-04 | 1.40E-05 | 0.36 | 4.46E-03 | 22  |
| 9.4  | 9.12 | 1.24E-02 | 3.77E-04 | 7.00E-06 | 0.01 | 4.46E-03 | 22  |
| 13.1 | 8.03 | 8.90E-03 | 3.50E-04 | 2.10E-05 | 0.71 | 7.24E-03 | 68  |
| 13.2 | 8.12 | 9.70E-03 | 3.50E-04 | 1.50E-05 | 0.71 | 7.17E-03 | 68  |
| 13.3 | 8.36 | 9.80E-03 | 3.76E-04 | 3.10E-05 | 0.71 | 6.92E-03 | 68  |
| 13.4 | 8.58 | 1.01E-02 | 3.53E-04 | 1.40E-05 | 0.71 | 6.57E-03 | 68  |
| 14.1 | 8.12 | 8.90E-03 | 3.39E-04 | 8.00E-06 | 0.02 | 7.05E-03 | 50  |
| 14.2 | 8.46 | 1.03E-02 | 3.21E-04 | 6.00E-06 | 0.02 | 6.55E-03 | 50  |
| 14.3 | 8.77 | 1.09E-02 | 3.68E-04 | 4.00E-06 | 0.02 | 5.76E-03 | 50  |
| 14.4 | 9.21 | 1.09E-02 | 3.86E-04 | 4.00E-06 | 0.02 | 4.07E-03 | 50  |
| 3.9  | 8.76 | 1.02E-02 | 1.08E-04 | 3.10E-05 | 0.71 | 2.31E-03 | 35  |
| 3.1  | 8.78 | 9.80E-03 | 7.30E-05 | 3.00E-05 | 0.71 | 1.53E-03 | 35  |
| 3.11 | 8.78 | 9.90E-03 | 7.70E-05 | 3.00E-05 | 0.71 | 1.38E-03 | 42  |
| 3.12 | 8.77 | 9.50E-03 | 6.20E-05 | 2.90E-05 | 0.71 | 1.07E-03 | 42  |
| 3.13 | 8.77 | 9.20E-03 | 6.20E-05 | 2.40E-05 | 0.71 | 7.70E-04 | 47  |
| 15.1 | 8.72 | 1.09E-02 | 3.22E-04 | 8.00E-06 | 0.02 | 5.91E-03 | 63  |
| 15.5 | 8.67 | 9.90E-03 | 3.31E-04 | 4.00E-06 | 0.01 | 3.03E-03 | 60  |
| 15.6 | 8.78 | 8.70E-03 | 2.32E-04 | 3.00E-06 | 0.01 | 1.43E-03 | 60  |
| 15.2 | 8.80 | 9.60E-03 | 1.66E-04 | 4.00E-06 | 0.01 | 7.10E-04 | 63  |
| 15.7 | 8.85 | 8.40E-03 | 1.19E-04 | 2.00E-06 | 0.01 | 4.40E-03 | 60  |
| 15.3 | 8.88 | 9.00E-03 | 9.30E-05 | 3.00E-06 | 0.01 | 3.24E-03 | 63  |
| 15.4 | 9.02 | 8.70E-03 | 4.30E-05 | 3.00E-06 | 0.01 | 9.70E-04 | 63  |
| 7.8  | 8.26 | 1.16E-03 | 9.80E-05 | 3.30E-05 | 0.70 | 8.0E-05  | 99  |
| 7.3  | 8.49 | 1.23E-03 | 1.02E-04 | 2.20E-05 | 0.70 | 8.74E-05 | 99  |
| 7.7  | 8.11 | 7.03E-03 | 9.00E-06 | 5.00E-06 | 0.70 | 7.71E-06 | 100 |
| 7.2  | 8.99 | 7.70E-03 | 1.00E-05 | 2.00E-06 | 0.70 | 8.57E-06 | 100 |
| 1.2  | 9.29 | 1.84E-02 | 1.09E-03 | 9.80E-05 | 0.70 | 9.31E-04 | 264 |
| 1.4  | 9.19 | 1.45E-02 | 7.69E-04 | 9.80E-05 | 0.70 | 6.59E-04 | 264 |
| 1.5  | 9.17 | 1.35E-02 | 5.76E-04 | 7.00E-05 | 0.70 | 4.94E-04 | 264 |
| 1.6  | 9.11 | 1.27E-02 | 4.52E-04 | 6.60E-05 | 0.70 | 3.87E-04 | 264 |
| 1.7  | 9.14 | 1.22E-02 | 4.00E-04 | 5.30E-05 | 0.70 | 3.43E-04 | 192 |
| 1.9  | 9.00 | 1.04E-02 | 1.77E-04 | 4.30E-05 | 0.70 | 1.52E-04 | 192 |
| 2.1  | 9.25 | 1.08E-02 | 2.85E-04 | 2.70E-05 | 0.70 | 2.44E-04 | 27  |
| 2.5  | 9.17 | 1.07E-02 | 2.43E-04 | 1.50E-05 | 0.70 | 2.08E-04 | 57  |
| 2.2  | 9.11 | 1.00E-02 | 1.99E-04 | 2.60E-05 | 0.70 | 1.71E-04 | 27  |
| 2.6  | 9.20 | 9.70E-03 | 1.49E-04 | 2.50E-05 | 0.70 | 1.28E-04 | 62  |
| 2.3  | 9.03 | 9.10E-03 | 9.60E-05 | 2.10E-05 | 0.70 | 8.23E-05 | 28  |
| 2.4  | 9.16 | 8.70E-03 | 9.10E-05 | 1.90E-05 | 0.70 | 7.80E-05 | 56  |
| 2.7  | 9.05 | 9.00E-03 | 6.30E-05 | 1.90E-05 | 0.70 | 5.40E-05 | 63  |
| 2.8  | 8.97 | 8.70E-03 | 5.40E-05 | 1.70E-05 | 0.70 | 4.63E-05 | 64  |

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**III. Comparison of peaks in Figure 2A, main paper, to peaks for known compounds.**

Table S-2. For each entry, the leading 3-digit number is the wavelength (nm) of an absorption peak and the number in parentheses is the molar extinction coefficient ( $M^{-1}cm^{-1}$ ).

| Anglin 1987<br>[(Fe <sub>2</sub> S <sub>2</sub> )(MoO <sub>2</sub> S <sub>2</sub> ) <sub>2</sub> ] <sup>4-</sup><br>H <sub>2</sub> O                      MeCN <sup>a</sup> . |            | McDonald 1980<br>[Fe(MoS <sub>4</sub> ) <sub>2</sub> ] <sup>3-</sup><br>MeCN <sup>a</sup> . | Coucouvani 1980<br>[Fe(MoS <sub>4</sub> ) <sub>2</sub> ] <sup>3-</sup><br>DMF <sup>a</sup> . | This Work<br>Figure 2A<br>H <sub>2</sub> O |
|-------------------------------------------------------------------------------------------------------------------------------------------------------------------------------|------------|---------------------------------------------------------------------------------------------|----------------------------------------------------------------------------------------------|--------------------------------------------|
| 290                                                                                                                                                                           | 290(23100) | no data                                                                                     | 290(21350)                                                                                   | 288(13150)                                 |
| 350                                                                                                                                                                           |            | 347(26920)                                                                                  | 348(25546)                                                                                   | 348(15750)                                 |
|                                                                                                                                                                               | 372(14600) | 372(sh16070)                                                                                | 375(sh)                                                                                      | 378(sh)                                    |
|                                                                                                                                                                               |            | 409(sh9150)                                                                                 | 409(9848)                                                                                    |                                            |
|                                                                                                                                                                               |            |                                                                                             |                                                                                              | 440(sh)                                    |
| 510                                                                                                                                                                           | 526(12200) | 507(18970)                                                                                  | 510(18254)                                                                                   | 506(7900)                                  |
|                                                                                                                                                                               |            | 577(7750)                                                                                   | 583(8360)                                                                                    | 574(5740)                                  |

<sup>a</sup>. MeCN=acetonitrile solvent; DMF=dimethylformamide solvent.

**IV. Trends in ΣFe and ΣMo in Well-Studied Modern Euxinic Basins.**

Table S-3 gives representative ΣFe and ΣMo concentrations in the deep waters of three well-studied euxinic basins. The basins are organized in the table so that ΣS<sup>-II</sup> in the deep waters increases downward. Going downward in the ΣFe<sub>aq</sub> column, the values are seen to flatten whereas in the ΣMo<sub>aq</sub> column the values pass through a minimum. These are the trends predicted in Figure 5 of the main paper.

Table S-3. Compositions of deep waters in three well-studied euxinic basins.\*

| Basin           | $\Sigma S^{-II}$ (mM) | $\Sigma Fe_{aq}$ ( $\mu M$ ) | $\Sigma Mo_{aq}$ ( $\mu M$ ) |
|-----------------|-----------------------|------------------------------|------------------------------|
| Cariaco         | 0.030                 | 0.300                        | 0.084                        |
| Black Sea       | 0.400                 | 0.038                        | 0.003                        |
| Framvaren Fjord | 6.000                 | 0.025                        | 0.018                        |

\* Data from: Helz et al. 2011; Emerson and Huested, 1991; Landing and Westerlund, 1988; Lewis and Landing, 1991, 1992; and Li et al. 2011.

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